

Hydro‘Carbon’ systems and what it can teach us about putting the ‘Carbon’ back, a CCS systems case study

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Introduction

Exploring for hydrocarbons is an incredibly complex pursuit. This is exemplified by Canadian exploration and exploitation budgets in 2023 – 40.5 billion CAD (StatCan, 2024). However, on another level it is very simple and can be broken down into the four basic play elements that are introduced in any university course that touches on hydrocarbon exploration. They are source, migration, reservoir, and trap/seal (SPE, 2024). There are strong similarities between the key elements of hydrocarbon systems and carbon capture and storage (CCS) systems. While not identical, the learnings from one can be employed to optimize the other. Here we explore how understanding the basics of a hydrocarbon system can help with CCS applied to sedimentary systems with a saline aquifer.

Source

Hydrocarbons are organic material (OM) composed entirely of hydrogen and carbon atoms that have been trapped within rock through sedimentary processes as kerogen, typically shales (i.e. source rocks).

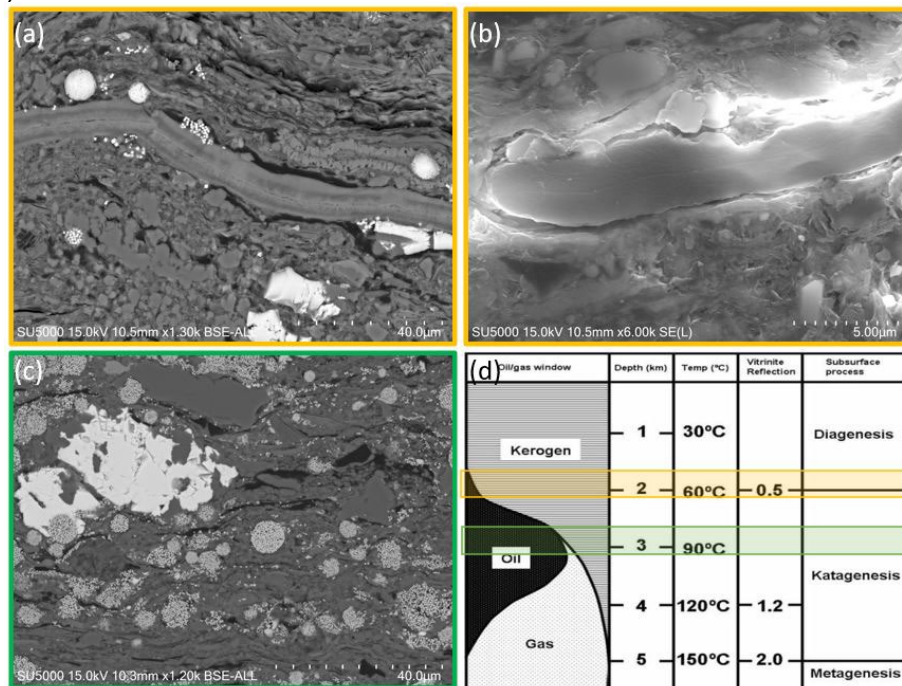


Figure 1 – (a) SEM image of well 16/8-3 S (Norwegian North Sea) highlighting immature kerogen lenses in black prior to cracking, with (b) a zoomed in image of a single kerogen lens surrounded by quartz and clay minerals. (c) SEM image of another well 15/9-4 (Norwegian North Sea) showing kerogen lenses that have partially cracked represented by black kerogen lenses with a grey ring around the outside edge. (d) Oil and gas window (from Tissot and Welte, 1984) with the oil/gas window that the samples from 16/8-3 S (yellow) and 15/9-4 (green) are currently in.

To be considered an economic source rock 2 wt.% total organic carbon (TOC) kerogen is an industry and academic standard that has been widely adopted (BGS, 2014). The conversion of the hydrocarbon from a solid (i.e. kerogen) to a fluid/gas occurs when the necessary temperature and pressure thresholds have been attained, and are referred to as the oil/gas window (Tissot and Welte, 1984). The reason for a window rather than a single point is that it is a combination of temperature, pressure, and time that allow for the conversion to take place (Gretnier and Curtis, 1982; Johnson et al., 2022b). Figure 1 shows a comparison of SEM images of shale containing kerogen that is immature and that have been partially cracked.

For CCS, the source could be considered analogous to the capturing of CO₂. This is done in a variety of ways. However, they can be broken down into direct air capture and industrial use capture (DNV, 2024). Direct air capture includes biological processes (e.g. macroalgae, microalgae) as well as technological (e.g. Climeworks). Industrial use capture can be further broken down into three categories, they are: post-combustion, pre-combustion, and oxy-fuel combustion (LSE, 2024). Once captured, the CO₂ can be transported to an injection site.

Migration

For hydrocarbon systems, migration describes the movement of fluid/gas from the source rock to a container (i.e. reservoir) with an effective lid (i.e. seal/trap). This process is simpler to outline than to understand. It requires expulsion from the low porosity (Φ), low permeability (K) source rock - the dynamics of which are still being debated and researched (e.g. Kobchenko et al., 2014; Johnson et al., 2022c).

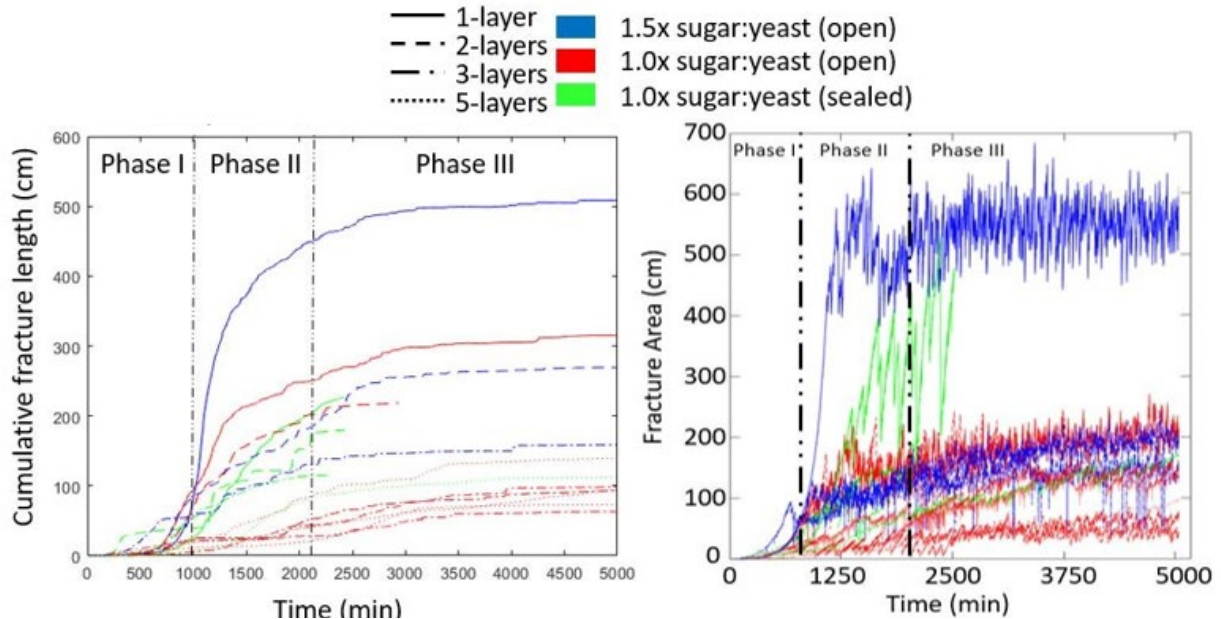


Figure 2 – Fracture network creation over time in an analogue model (left) showing Phase I (diffusion and fracture creation), Phase II (fracture network interconnection), and Phase III (fluid expulsion). The fractured area stabilizes in Phase II during final fracture interconnection, and then varies within a given range in Phase III for each experiment as the fractures open and close during fluid expulsion.

One model for expulsion is explored by Johnson et al., (2022c) wherein the process is originally driven by diffusion. However, low Φ - K relationships result in saturation and subsequently fracture

propagation (Figure 2) driven by the largest kerogen lenses (Johnson et al., 2022b). With appropriate kerogen distribution within the shale this will create a fracture network that will open and close as pressure rises and falls (Johnson et al., 2022b; Johnson et al., 2022c), allowing the periodic release of hydrocarbons (Figure 2).

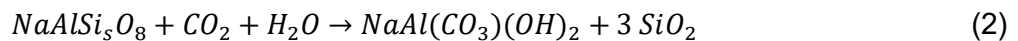
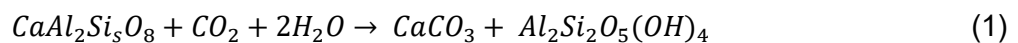
Movement of hydrocarbons within the subsurface are driven by buoyancy, gravitational, and capillary forces. At the most basic level, the pathways are determined by Φ -K relationships via geological composition and structural elements (e.g. fractures, faults) (Wangen, 1995; Manzocchi et al., 2010). Hydrocarbons are geochemically reasonably stable, drastically reducing the potential interactions during migration (Price, 1993).

Initial projects for CO₂ storage assumed little to no migration would occur (Ringrose et al., 2013), however pilot studies for CCS has shown this to be a faulty assumption (e.g. In-Salah, Snøhvit). Therefore it is necessary to consider not only primary, but secondary seals in addition to possible migration pathways (Johnson et al., 2022a). Unlike hydrocarbons, CO₂ can be incredibly reactive when in the presence of the appropriate mineralogy. Busch et al., (2016) highlighted that CO₂ can be especially reactive with smectite resulting in volumes changes associated with fracture creation. While the injection of CO₂ does carry risk with it including seismic/aseismic fault slip, seal/trap integrity damage, and uplift/subsidence (Rutqvist, 2012). However, in the case of CO₂ escape from the primary reservoir the worst case scenario for CO₂ migration is that it is re-released into the atmosphere after a tortuous journey through the subsurface. This does however highlight the need for surface monitoring at all CCS wellsites.

Reservoir

There is a long history within hydrocarbon exploration of targeting primarily two reservoir types (i.e. sandstones, carbonates). Worldwide, the ratio of sandstone to carbonate reservoirs are 3:2 (Bjørlykke and Jahren, 2010) while the distribution of total hydrocarbons in sandstone reservoirs is closer to ~50% (SLB, 2024).

Similarly, both sandstone and carbonate reservoirs have been considered and utilized for CO₂ storage pilot projects (i.e. <1 Mt /yr). Since CO₂ is geochemically reactive, larger projects have favoured sandstone reservoirs preferably composed dominantly by quartz (Figure 3). Deviations from this encourage geochemical reactions including, but not limited to (Hangx and Spiers, 2009):



The potential problem with numerous geochemical reactions is that it greatly increases uncertainty, including increased potential for complications as a result of injection as outlined by Rutqvist (2012). For example a feldspar rich sandstone could see a reaction similar to Equation 1, in which anorthite goes to CaCO₃ and kaolinite in the presence of CO₂ and water. CaCO₃ will likely react further (see Equation 3) while the presence of kaolinite will likely clog the pore space.

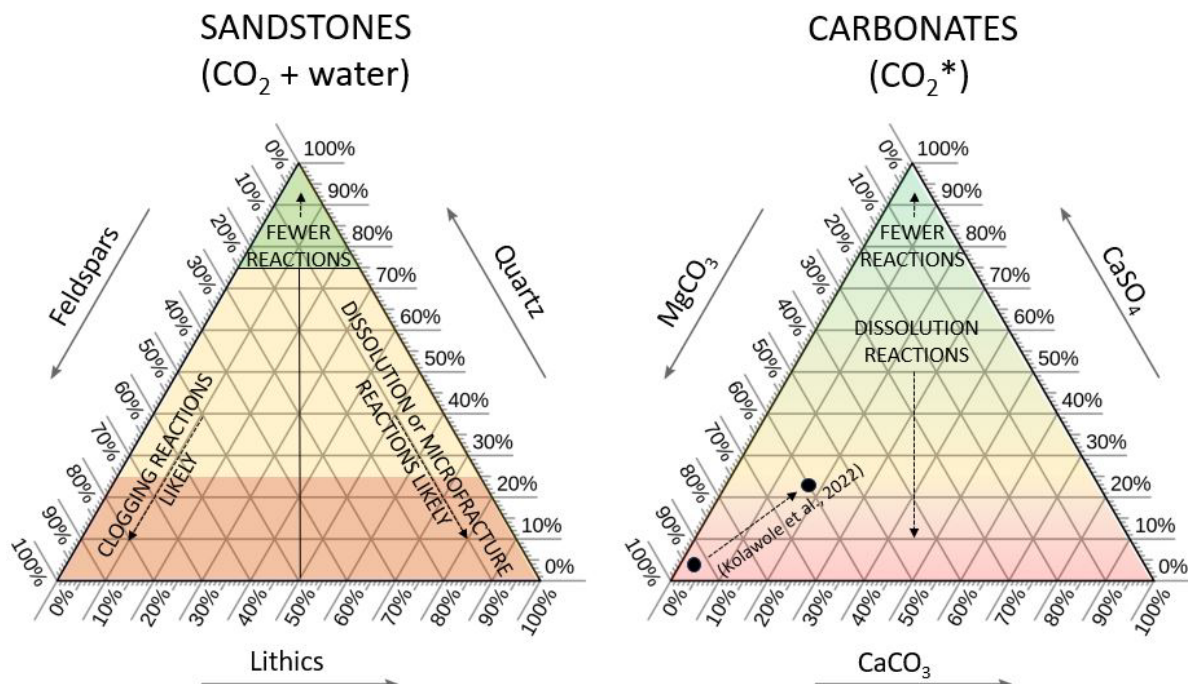
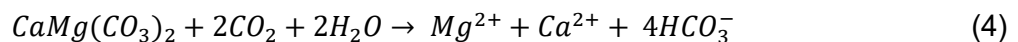


Figure 3 – Shows less reactive regions of mineralogical content of a sandstone (left) and carbonate (right) reservoir given current understanding of geochemical reactions between CO₂ and minerals commonly present within these reservoir types. On the sandstone ternary diagram the presence of water is also considered. On the carbonate ternary diagram, the impact of introducing *Sporosarcina pasteurii* on the mineralogical content is highlighted based on work by Kolawole et al., (2022).

Carbonate reservoirs have been considered especially problematic as most carbonate minerals readily react with CO₂. For example (Campbell et al., 2022):



It has been established that these alterations can lead to dissolution within the reservoir, encouraging mechanical weakness (Seyyedi et al., 2020). However, many sedimentary basins do not have an abundance of clean quartz reservoirs and may at least partially rely upon the use of carbonate reservoirs for their CCS opportunities (e.g. Western Canadian Sedimentary Basin, Central Arabic Basin). To this end, early research has shown that there are potential pathways forward (Kolawole et al., 2022). For example, injecting a preparation fluid into a carbonate reservoir can significantly alter the geochemistry in such a way that it would be more ready to receive CO₂ (Figure 3).

In the case of all reservoirs for CCS, it is important to consider when CO₂ becomes supercritical. At approximately 800m depth, CO₂ has been shown to become supercritical, although this can vary somewhat depending on local PT regime. When CO₂ becomes supercritical the amount of CO₂ that you can fit into a given space is both significantly improved, as well optimized. Roughly

800m is considered optimal since the further improvement for storage capacity between approximately 800m and 3000m depth is 0.05% compared to the compression that has already occurred during the first 800m (DOE, 2024). Wang et al., (2018) has shown that cooler basins will result in greater CO₂ density as pressure and depth increase.

Seal/Trap

When considering potential seals/traps for hydrocarbons there are a variety of geometries that could potentially work, however the most commonly used and the safest are 4-way closures and 3-way closures (SPE, 2024). 4-way closures depend entirely on the shape of the seal in combination with an appropriate Φ -K relationship (i.e. low Φ , low K). 3-way closures are more complex in that they further depend on a large-scale structural element (i.e. fault) and the compositional relationships across them (Yielding et al., 2010). The classic relationship shows the importance of these compositional relationships. However, it goes on to explain that this can be further complicated by the sealing or non-sealing nature of the fault gouge itself (Yielding et al., 2010; Vrolijk et al., 2016). This is another field of research within which there is an active and ongoing debate (Smith, 1966; Yielding et al., 2010; Vrolijk et al., 2016).

For CO₂, we have already touched on the potential for geochemical reactions and how that might adversely affect a potential seal. Therefore, a focus will be applied on understanding the long-term advantages and disadvantages of these potential seal/trap types. Since 3-way closures depend on a sealing-fault, and subsurface systems are dynamic especially when interacted with, they present a greater risk than 4-way closures. However, most if not all basins have undergone some major structural changes via tectonic forces resulting in the need to consider more complex geometries frequently (e.g. WCSB, Norwegian Continental Shelf).

Due to the comparable size and attributes of CH₄ and CO₂, it could be argued that hydrocarbon reservoirs that have had supercritical CH₄ have 'tested capacity', while saline aquifers do not. Indeed, where this is the case and space has been made available these areas should be primary targets for permanent CO₂ storage. However, due to the limited amount of space available in all hydrocarbon reservoirs, not just the ones that had supercritical CO₂, it will be necessary to utilize saline aquifers (Anthonsen et al., 2013). Nonetheless, understanding the characteristics of traps/seals that held supercritical CH₄ may provide the greatest learnings for future CO₂ reservoirs.

Conclusions

Here we have utilized the same elements of a hydrocarbon system to look at the components of CO₂ systems, and the different kinds of complexity that are associated with a change in fluid in combination with a change in direction of fluid flow. The development of larger scale CO₂ projects are and should develop from the simplest combination of these elements (e.g. Sleipner, Quest). However, the sheer size of the problem demands innovation that will allow more complex CO₂ systems to be harnessed. Complex sourcing (e.g. DACs), complex migration (multiple seals), complex reservoirs (e.g. carbonates, arkose, lithic), and complex seal/trap sequences should be viewed as necessary scientific challenges in order to dispose of the anthropogenically created waste product (i.e. CO₂).

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